REPORT ON 2017 GEOPHYSICAL SURVEYS
GOODMAN PROPERTY

Yukon, Canada

136˚ 03’ W
63˚ 57’ N

October 31, 2017

Prepared for:
GENERIC GOLD CORPORATION

Prepared by:

AURORA GEOSCIENCES

MODELLING REPORT
2017 GEOPHYSICS (IP SURVEY)
GOODMAN PROPERTY, YUKON

Effective Date: October 31, 2017

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1 SUMMARY

During September and October 2017, a direct-current resistivity and induced polarization (DCIP) survey was conducted on Generic Gold Corporation's Goodman Property. Six profiles were surveyed, comprising a total of 10.6 line-km of DCIP survey. This report describes inversion modelling of the collected DCIP data. Modelling is performed using DCIP2D software, developed at UBC-GIF, which provides a two dimensional distribution of sub-surface physical properties that are consistent with results measured by the geophysical survey.

Details of survey parameters and equipment used can be found in the field report "2017 Goodman IP Field Report" dated October 15th, 2017. DCIP data were acquired on parallel survey lines spaced 400 m apart utilizing a pole-dipole electrode configuration. Ten dipoles were measured with potential electrodes spaced 25 m apart. Electrode locations were recorded with non-differential GPS receivers. Station elevations were extracted from Canadian Digital Elevation Data for 1:50,000 NTS sheets provided by GeoGratis Client Services.

The IP survey covers an extensive package of Yusezyu Formation sedimentary rocks. Felsic intrusive rocks are interpreted from an airborne magnetic survey to underlie the northwestern extents of the grid. A deposit model for the Property would conform to an intrusion related gold model and gold mineralization would be accompanied by disseminated sulphide mineralization occurring as veins, stringers or stockwork zones. Sulphide mineralization can be detected from the surface using the DCIP technique which measures the elevated subsurface chargeability caused by the presence of such mineralization. The survey also provides resistivity information which can be used as an exploration tool in this particular geological setting.

Inversion modelling of the IP survey results identified 18 discrete chargeable sources which may be mapping the locations of disseminated sulphide mineralization underlying the survey area. Resistivity modelling indicates that the chargeability bodies are associated with either an increase or decrease in host rock resistivity. An anomaly summary map (Figure 1) shows the locations of the identified chargeability and resistivity anomalies in relation to historical soil geochemistry and airborne magnetic survey results.

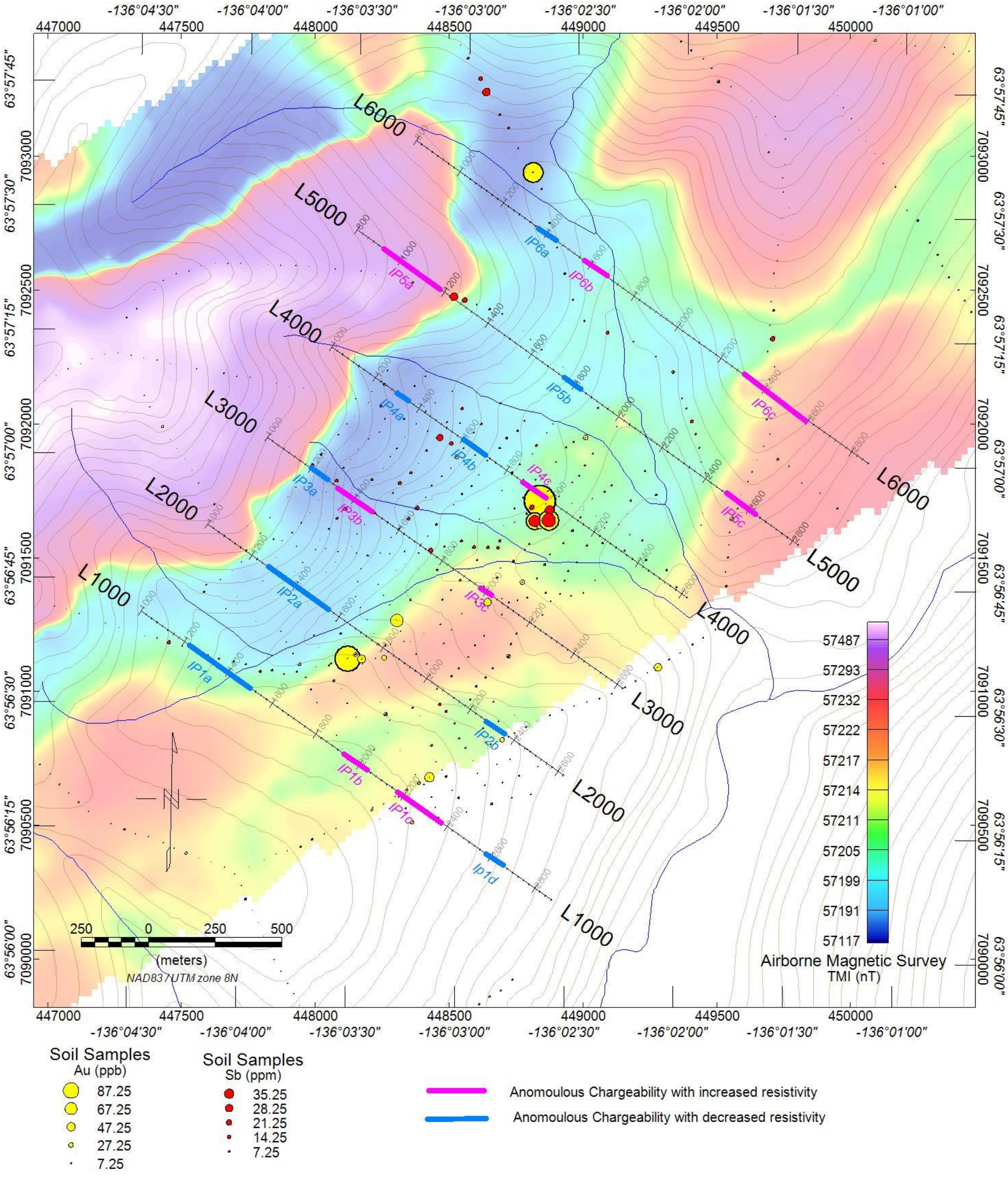

Figure 1 : Anomaly overview

2 DCIP INVERSION MODELLING

Pseudosections of DCIP data estimate the depth and location of features based on the location and separation of transmitting and receiving electrodes. Inverting the DC resistivity and IP data produce 2D models of recovered resistivity and chargeability with true depths. The 2D models are distributions of rock properties with varying chargeability and electrical resistivity which produce responses that match the observed responses within the limits of instrument and survey error. There is no unique distribution of chargeability and resistivity which will create the observed responses; rather there is a range which can be narrowed by incorporating geological information to restrict the possible solutions.

The DCIP data are inversion modelled using the software DCIP2D, developed at the University of British Columbia's Geophysical Inversion Facility. Quality control was performed prior to export; any reading with an irregular decay curve or suspect primary voltage is excluded. Repeat readings are averaged. The core survey area covering all receiver electrodes are discretized by 12.5 m cells horizontally and 6.25 m cells vertically. Additional horizontal and vertical padding cells, each with a cell expansion ratio of 1.5 are added to the mesh prior to inversion; these cells are stripped from all products attached to this report. Topography is accounted for using a 10 m digital terrain model generated by gridding topographic map elevation data from NTS sheet (1: 50 000 115P/16).

Normalized voltage potentials calculated from the measured voltages and transmitting currents are first inverted to recover a conductivity model required for the chargeability inversion. Voltage potentials for the resistivity calculations are assigned an error of 5% plus 0.001 S/m, chargeability errors assigned 5% plus a minimum of 0.5 mV/V. The limited amount of available geological information for the survey area does not allow the application of model constraints and default inversion parameters are chosen to generate a model equally smooth in horizontal and vertical directions. Survey lines are individually modelled and both resistivity and chargeability 2D inversions are run a second time using greater reference models to provide an estimate of the survey array's sensitivity and depth of investigation.

Satisfactory model solutions which reproduce the observed data were obtained for all of the survey lines.

3 RESULTS

The results of the IP inversions are presented as depth sections of chargeability and resistivity and are included in Appendix 1. The sections are plotted looking northeast and without vertical exaggeration. Coordinate axes are labelled with UTM coordinates and elevations, all in metres. Grid coordinates used by the DCIP survey crew during data acquisition are annotated for reference along the bottom of each section and along the section trace plan view included on each section. The section trace is sketched over a historical airborne magnetic survey field image and includes watercourses and elevation contour lines. A common colour scheme is assigned to both the chargeability and resistivity model images to allow a line to line comparison of modelled values.

An examination of the measured data plotted as pseudosections shows a general increase in chargeability with depth over the entire survey area. This background response is interrupted with discrete, well-formed anomalies extending beyond the survey's depth of investigation. Resistivity values are more variable both with depth and down line distance. The observations suggest that the survey area is almost entirely overburden covered, and that the chargeability anomalies are likely due to bedrock sources while resistivity anomalies may be resulting from either bedrock or overburden sources.

Model results show resistivity values ranging between 50 to 13000 ohm-m with a mean value of approximately 750 ohm-m, and chargeability values ranging from 0 to 50 mV/V with a mean value of approximately 15 mV/V over all survey lines. Statistics for the individual survey lines are shown in Table 1.

Table 1 : Model statistics

Line	Range of Model Chargeability (mV/V)	Mean Chargeability (mV/V)	Range of Model Resistivity (Ohm-m)	Mean Resistivity (Ohm-m)
1	0-45	14	50-11000	680
2	0-30	14	70-5000	600
3	0-50	18	75-13000	700
4	0-30	15	100-6000	600
5	0-45	20	70-9000	1000
6	0-37	16	70-7000	900

Based on these values the modelled chargeability sources can be categorized as strong (> 30 mV/V), moderate (20 to 30 mV/V) and weak (15 – 20 mV/V). The large range in resistivity values reflect both overburden and bedrock contributions. Decreased bedrock resistivity is estimated to be indicated by values of less than 400 ohm-m and increased resistivity by values greater than 900 ohm-m.

The identification of anomalous chargeability and resistivity is limited to the examination of deep seated features extending beyond the survey's depth of investigation. Shallow and / or flat lying elements likely caused by surficial sources are not considered in this interpretation.

The locations of identified chargeability anomalies are listed in Table 2 and sketched on the accompanying model sections. A brief description for each of the anomalies follows.

3.1 Line 1

3.1.1 Anomaly IP1a

This broad, moderately chargeable source is located between grid coordinates 1200 – 1600. The body's orientation is sub-horizontal and associated with moderate to low resistivity. It appears to be buried approximately 110 m below surface.

3.1.2 Anomaly IP1b

A deep seated weakly chargeable source is modelled to lie between grid coordinates 1900-2000 and is associated with moderately low resistivity. The depth of emplacement is roughly 125 m.

3.1.3 Anomaly IP1c

A strongly chargeable source is located between grid coordinates 2100 – 2400. The depth below surface to the top of the body is approximately 50 m. The body is associated with moderately high resistivity.

3.1.4 Anomaly IP1d

A modelled moderate chargeable source is centered on grid coordinate 2600 at a depth of roughly 90 m. The anomaly is located within an area of low resistivity.

3.2 Line 2

3.2.1 Anomaly IP2a

This broad, moderately chargeable anomaly is located between grid coordinates 1200 – 1500 and is associated closely with an area of low resistivity. The depth to the top of the body is approximately 100 m. The orientation of the body appears to be sub-horizontal.

3.2.2 Anomaly IP2b

A weakly chargeable source is modelled to lie between grid coordinates 2250-2350 at roughly 125 m depth. The body is coincident with a sub vertical dipping resistivity low.

3.3 Line 3

3.3.1 Anomaly IP3a

A strongly chargeable source centered on grid coordinate 1225 is modelled to extend within 35 m depth from surface. The source is coincident with a resistivity low.

3.3.2 Anomaly IP3b

A broad, strongly chargeable feature is located between grid coordinates 1325-1600. The depth to the top of the body increase from approximately 25 m near station 1300 to greater than 100 m near station 1600. The anomaly straddles a contact between a resistivity high and low.

3.3.3 Anomaly IP3c

The top of a strongly chargeable body is modelled to lie between grid coordinates 1850-2000. The body is associated with moderate resistivity and is overlain by approximately 135 m of low resistivity material.

3.4 Line 4

3.4.1 Anomaly IP4a

A loosely defined zone of moderate chargeability approximately 75 m wide is centered on grid coordinate 1325. The zone is coincident with a distinct sub vertical resistivity low feature which extends to the near surface.

3.4.2 Anomaly IP4b

An area of moderate chargeability approximately 150 m wide and occupying an area of moderate resistivity is located between grid coordinates 1600-1750. The depth to the top of the zone is roughly 90 m.

3.4.3 Anomaly IP4c

A 100 m wide zone of moderate chargeability is modelled between grid coordinates 1875 – 1975. The top of the zone lies at approximately 75 m below surface. A subdued resistivity high is associated with this anomaly.

3.5 Line 5

3.5.1 Anomaly IP5a

A strongly chargeable zone 250 m in width is modelled between grid coordinates 925-1175 at a depth of approximately 75 m. The feature occupies an area of elevated resistivity.

3.5.2 Anomaly IP5b

The top of a strongly chargeable zone is modelled to lie within 40 m of the surface, centered at grid coordinate 1800. The zone is coincident with a resistivity low anomaly.

3.5.3 Anomaly IP5c

A moderate chargeability source approximately 75 m wide is located between grid coordinates 2500-2575 at a depth of roughly 75 m below surface. The entire body is modelled to lie within an area of high resistivity.

3.6 Line 6

3.6.1 Anomaly IP6a

A moderately chargeable zone approximately 75 m wide and associated with low resistivity is modelled at grid coordinates 1400. The top to this body is approximately 75 m below surface.

3.6.2 Anomaly IP6b

A moderately chargeable zone approximately 75 m wide and at a depth of 75 m is modelled at grid coordinate 1600. This zone is related to moderately high resistivity.

3.6.3 Anomaly IP6c

A broad zone of moderate to strong chargeability is modelled to lie between grid coordinates 2300-2600. The zone appears to comprise two lobes of strongly chargeable and resistive material separated by a narrow sub vertical oriented resistivity low.

Table 2 : Anomaly locations. Coordinates define the centre of the anomaly.

Anomaly ID	UTM Easting Coordinate (NAD 83 zone 8)	UTM Northing Coordinate (NAD83 zone 8)	Depth (below surface to top) (m)
IP1a	447640	7091100	80
Ip1b	448140	7090740	135
IP1c	448383	7090570	50
IP1d	448670	7090370	90

IP2a	447925	7091395	70
IP2b	448680	7090860	115
IP3a	448015	7091815	25
IP3b	448190	7091690	25
IP3c	448580	7091415	125
IP4a	448320	7092105	80
IP4b	448585	7091920	85
IP4c	448800	7091765	60
IP5a	448350	7092585	50
Ip5b	448950	7092160	50
IP5c	449575	7091710	100
IP6a	448850	7092720	75
IP6b	449025	7092595	85
IP6c	449650	7092150	50

4 CONCLUSIONS AND RECOMMENDATIONS

The following conclusions are based on the DCIP modelling and interpretation described in this report.

- Bedrock hosted chargeability sources occur in several locations throughout the survey area.
- The chargeability sources are modelled to lie beneath 25 -125 metres of cover
- The chargeability sources can be divided into those associated with increased and those associated with decreased rock resistivity.

The chargeability response from rocks containing disseminated sulphides depends largely on several factors including but not limited to the type of mineralization, the volume content of sulphide minerals, the absolute size and shape of the sulphide grains, and the geometry of the sulphide body and its location relative to the measuring array. In addition to the chargeability response from sulphides, rocks containing magnetite, graphite, clay minerals or variably altered rocks may also produce chargeability responses of varying amplitudes.

The DCIP survey on the Goodman Property outlines areas of increased chargeability which may be the result of disseminated sulphide mineralization but further work would be required to fully define the source of the anomalies and to determine if economic mineralization is present.

Further work would include drill testing of the chargeability sources described in this report. Initial holes should include chargeable sources associated with both increased and decreased resistivity. Should the drilling successfully intercept a mineralized structure a geophysical signature could then be established that would be used to guide further exploration.

Respectfully Submitted,

Aurora Geosciences Ltd.

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Appendix 1 – Model Sections